

Thallium in deep-seated crustal rocks

G. PAOLO SIGHINOLFI

Istituto di Mineralogia, Università di Modena, Modena, Italy

ADELAIDE M. SANTOS

Instituto de Geociencias, Universidade Fed. da Bahia, Salvador, Brazil

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Abstract—A Precambrian granulite terrain of the Brazilian shield presents an average Tl content of 350 ppb, very near to Shaw's figure for the lower continental crust. Thallium, as in other rock types, shows close coherence with Rb and K. The mean Rb/Tl value (193 for 53 samples) falls in the range of the 'normal' crustal values, while the K/Tl values (145,000 average) are higher than values for igneous rocks, as are the values for K/Rb. The results support the authors' hypothesis that granulitic rocks have undergone partial melting with contemporaneous depletion of Rb and Tl in relation to K.

INTRODUCTION

VERY SCANTY data are available on thallium distribution in metamorphic formations, particularly in high-grade metamorphic rocks and nothing is known about thallium behaviour in sequences of increasing grade of metamorphism. Conclusions of some investigations (see the review by DE ALBUQUERQUE and SHAW, 1972) show that the thallium content of various metamorphic rocks depends mainly on the mineralogical composition, and that granulitic rocks may be depleted in thallium in relation to other rocks. Granulites are normally considered (see SHAW, 1972) as deep-seated crustal materials, and therefore trends followed by elements in granulites may be considered representative trends of the lower continental crust. The present paper describes thallium behaviour in a granulite terrain in the Precambrian Brazilian shield on which a number of chemical studies have already been published (SIGHINOLFI, 1970, 1971, 1972).

Thallium was determined in 62 samples by flameless A.A. spectroscopy using the Perkin-Elmer heated graphite atomizer HGA-70 and Model 303 spectrophotometer. Thallium was previously extracted in an ether phase following the procedure used by FRATTA (1973). The detailed method here used is described by SIGHINOLFI (1973). Sample aliquots of 0.2 g were used and the detection limit can be placed at about 20 ppb. Precision of the bulk analysis depends on the Tl levels in the sample. In the range 20–50 ppb an error of 30–50 per cent must be considered; for Tl content above 50 ppb, precision becomes good (error < 10 per cent).

Accuracy checks were carried out on USGS standard rocks. Duplicate determinations on each standard give the following values (ppb): W-1 105–110, G-2 950–1180, AGV-1 535–590, GSP-1 1520–1550 and BCR-1 280–300.

ANALYTICAL RESULTS

Table 1 reports thallium data and includes values of the Rb/Tl and K/Tl ratios. Eight samples are characterized by Tl levels below the detection limit (20 ppb) and in computing the average Tl value, they were arbitrarily considered to contain

Table 1. Thallium content and Rb/Tl and K/Tl values

Sample	Tl (ppb)	Rb/Tl	K/ Tl × 1000	Sample	Tl (ppb)	Rb/Tl	K/ Tl × 1000
II 1	<20	—	—	II 47	150	507	207
II 6	30	300	123	II 50	990	124	40
II 7	20	250	155	II 52	270	341	158
II 8	120	142	77	II 53	<20	—	—
II 9	270	130	78	II 54	100	830	362
II 10	70	814	366	II 55	>10000	—	—
II 12	120	108	67	II 56	220	114	84
II 14	20	280	120	II 57	940	97	39
II 17	40	140	70	II 58	1620	116	23
II 18	940	87	32	II 59	520	104	40
II 19	30	253	160	II 60	<20	—	—
II 24	640	91	39	II 61	<20	—	—
II 25	50	200	122	II 62	130	223	145
II 26	180	139	83	II 63	40	225	175
II 27	110	118	111	II 64	40	250	122
II 28	40	150	145	II 66	<20	—	—
II 29	<20	—	—	II 67	<20	—	—
II 30	<20	—	—	II 68	280	179	90
II 31	415	106	47	II 69	220	159	95
II 32	970	114	30	II 70	290	148	57
II 33	480	108	30	II 71	140	143	99
II 34	2640	80	21	II 72	200	210	72
II 36	80	462	275	II 73	360	167	63
II 38	350	97	70	II 74	480	119	42
II 39	530	149	57	II 76	30	233	333
II 40	990	93	23	II 77	1140	91	32
II 41	790	115	36	II 78	120	83	58
II 42	1410	83	28	II 79	320	87	62
II 43	80	187	102	II 82	70	143	194
II 44	190	226	135	II 83	390	87	57
II 45	380	216	68	Mean	350*	193†	145†
II 46	250	228	108				

* Number of samples = 61 (see text).

† Number of samples = 53.

10 ppb Tl. This procedure cannot influence sensibly the calculated thallium average for the whole terrain, and these eight samples were not considered in computing the averages for Rb/Tl and K/Tl ratios.

Sample II 55, characterized by a very high Tl content (more than 10 ppm), was also excluded from all the calculations, in order not to falsify the figures for the whole terrain. This assumption is justified because this thallium anomaly seems to be due to a localized process which later affected the terrain. In fact, sample II 55, even if mineralogically true granulite, has an abnormally high microcline content, sometimes concentrated in the form of recrystallized veins, and later hydrothermally deposited K-felspar (particularly of amazonitic variety) is known to be very rich in thallium (SITNIN, 1960).

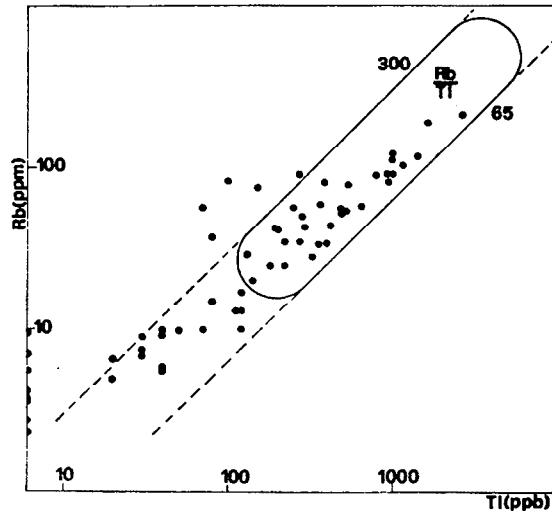


Fig. 1. Rb-Tl relationship. Area between dashed lines: terrestrial Rb/Tl values. Area outlined: composition ranges for igneous rocks (from DE ALBUQUERQUE and SHAW, 1972).

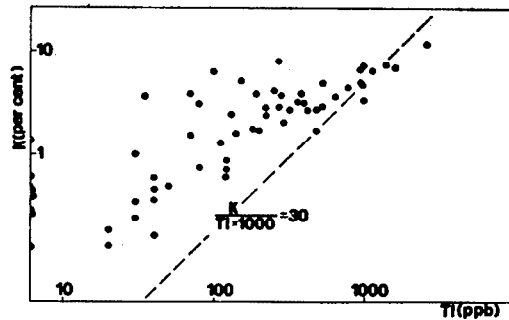


Fig. 2. K-Tl relationship. $K/Tl \times 1000 = 30$: average value for igneous rocks (DE ALBUQUERQUE and SHAW, 1972).

Tl-Rb-K RELATIONSHIPS AND IMPLICATIONS ON DEEP CRUSTAL PROCESSES

The geochemistry of thallium in the granulites analysed, as in igneous rocks, is dominated by the close coherence with rubidium and potassium (Figs. 1 and 2). The average thallium content in the whole terrain (350 ppb) agrees with the figure, deduced by extrapolations by SHAW (1972), for the lower continental crust (400 ppb). The Rb/Tl values fall for the most part within the range of 'igneous' values (Fig. 1) and the mean value (193) is slightly higher than the average for igneous rocks (about 150, following DE ALBUQUERQUE and SHAW, 1972). The K/Tl values, on the contrary, are markedly higher than most 'igneous' values, and the mean value is two or three times greater than some proposed averages for both igneous rocks and oceanic tholeiites (see Table 3). From the data of Table 2, thallium appears to be more strongly correlated with rubidium than with potassium. Furthermore,

Table 2. Correlation coefficients and some linear regression data relative to thallium, rubidium and potassium

x	y	\bar{x}	\bar{y}	n	r_{xy}	t
Tl	Rb	402	51	53	0.90	15.2
Tl	K	402	2.05	53	0.76	8.3

ppb Tl = 9.946 (ppm Rb) - 103.9
ppm Rb = 0.082 (ppb Tl) + 17.7
per cent K = 0.01 (ppb Tl) + 1.3

n = degrees of freedom; $t_{1\%}^2$ for 51 ($n - 2$) degrees of freedom = 2.68.

Table 3. Thallium in various Earth's crust units

	Thallium (ppb)	Rb/Tl	K/ Tl $\times 1000$	
Whole granulite terrain				This work
number of samples	61	53	53	
range	20-2640	80-830	21-362	
average	350	193	145	
Igneous rocks				DE ALBUQUERQUE and SHAW, 1972
range		19-590	4-112	
average		150	30	
Oceanic tholeiites				DE ALBUQUERQUE <i>et al.</i> , 1972
number of samples	5	5	5	
range	23-63			
average	42	162	96	
Whole Earth's crust				DE ALBUQUERQUE and SHAW, 1972
average	800			
Upper continental crust				SHAW, 1972
average	800			
Lower continental crust				SHAW, 1972
average	(400)			
Oceanic crust				SHAW, 1972
average	50			

regression analysis data show that the Rb/Tl ratio remains roughly invariant at different Rb levels while the K/Tl ratio decreases strongly with increasing K content. This makes the K/Tl trend similar to the K/Rb trend recognized in the same granulite terrain (SIGHINOLFI, 1970, 1971) just as in other granulites. In recent years, abundant evidence has accumulated indicating that high-grade granulite facies terrains are depleted in some lithophile elements including K, Rb, U and Th (see a review in SIGHINOLFI, 1971): from the results of this work conclusions are that thallium also was depleted from granulites at least to the same extent as of rubidium. Mechanisms which have been suggested as the cause of the depletion of elements from deep rocks include anatexic melting, with or without migrating melt phase, and progressive dehydration reactions but the effectiveness of both processes is still far from being known. Thallium data in granulites can contribute to understanding ion transport mechanisms in the deep crust. In igneous processes thallium and rubidium enter the same phases (K-minerals) and DUPUY *et al.* (1973) show that

their partition coefficients are about the same in most of the phases. In the course of partial melting, therefore, the two elements may enter the melted fraction in substantially the same proportions and the Rb/Tl ratio would not vary sensibly either in the unmelted residue or in the melt. This would be true whatever may have been the evolutionary course of the partial melting process (i.e. melting and almost complete recrystallization *in situ* or melting followed by migration of the large part of the K-rich melt).

The other mechanism, i.e. dehydration reactions, involves the formation of a fluid phase in which both rubidium and thallium are known to enter and migrate easily, for example, as halides or fluorides (VOSKRESENSKAYA, 1959), but no precise information is yet available on the relative behaviour of thallium and rubidium in a gaseous or aqueous phase in equilibrium with a rock at variable P - T conditions. In the opinion of MAROWSKY and WEDEPOHL (1971), thallium is prevented from major fluid transport by fixation in potassium minerals. VOSKRESENSKAYA and FEL'DMAN (1964) stated that thallium and rubidium become partially separated during alkali metasomatism and thallium is more mobile than rubidium.

On the basis of these considerations, some results of this paper (Rb and Tl depletion rate in respect to K, no apparent fractionation between Rb and Tl in granulite residuals) seem to indicate that partial melting alone could have established the observed Tl-Rb-K trends, even though dehydration reaction effects cannot be excluded completely. Whatever the mechanism of depletion (although more and more evidence suggests that melting plays the major role), according to recent studies by LEWIS and SPOONER (1973), the major depletion of lithophile elements and therefore also of thallium from deep-seated rocks may have occurred in the early stages of crustal evolution rather than by repeated removals through time.

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